

Impact of cross-reef water fluxes on lagoon dynamics : a simple parameterization for coral lagoon circulation model, with application to the Ouano lagoon, New Caledonia

Cristele Chevalier · Damien Sous · Jean-Luc Devenon ·
Marc Pagano · Gilles Rougier · Jean Blanchot

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Abstract This manuscript presents a combined experimental and numerical study of the impact of cross-reef fluxes on coral reef lagoon dynamics. The selected field site is the Ouano lagoon (New Caledonia island, France) in the South Western Pacific Ocean. Measurements of wave transformation above the reef and current profiles through passages and reef openings have been carried out during a three months survey. Data analysis reveals the preponderant roles played by both tides and waves on the lagoon dynamics. Based on field data, a simple parameterization of cross-reef fluxes is implemented in a coastal lagoon circulation model and a satisfactory agreement is found between parameterized model and field results. The model is thus used as a numerical experimental tool in order to analyse the cross-reef flows possible influence on a narrow lagoon dynamics. The results highlight the importance of cross-reef fluxes induced by wave breaking over the reef barrier on the whole lagoon circulation and water properties.

Keywords Coral reef lagoon · Lagoon hydrodynamics · Cross-reef fluxes · Circulation model · Channel lagoon

1 Introduction

Mostly present in intertropical regions, coral reef lagoon systems provide a unique habitat for a wide variety of living organisms. The lagoon biodiversity is dependent on the health of the whole food chain, in particular on plankton composition and abundance [1], and is then largely reliant on the water quality of the whole system. As a living organism, the reef itself (health, size and growth) is also particularly sensitive to the quality of surrounding marine waters, the aeration and mixing processes and the richness in nutrients [2, 3].

The quality of the reef-lagoon system waters and their bio-chemical properties are strongly dependent on circulation patterns and on water renewal time, which are mainly determined by flows through passages and reef openings and over the coral reef [4,5]. These exchange flows between open ocean and lagoon are controlled by a set of external parameters, including tides, waves, winds, rainfalls and river discharges which spatio-temporal variations eventually rule the lagoon circulation. The relative contributions of reef and passages to the ocean-lagoon exchanges are also determined by geometrical considerations on bathymetry as the ratio between vertical cross-sections above the immersed reef flat and through the passages. Whereas the flow through passage and reef opening is weakly dependent on the water level, the flow above the reef barrier may be strongly depend on tidal fluctuations. At high tide, flows across the reef can be at least as great as flows through the passages and may even dominate in the extreme case of a shallow lagoon compared to the tidal range, see e.g. the Bamburi lagoon, Kenya, which is exposed to tidal oscillations up to 3m and where the maximal depth is less than 10m and even inferior to 3m over a large part of the lagoon [6]. Conversely, flows across the reef become almost negligible at low tide when the reef is slightly submerged or totally exposed.

A significant amount of research works have been dedicated to the understanding of small-scale hydrodynamics in and over coral reefs thanks to field measurements [7–14], laboratory experiments [15–17], analytical approaches [18,2] or modeling studies [19], see Monismith [20] for a review. Cross-reef fluxes

have been observed to be primarily driven by the combined effects of tidal oscillations and, when significant, wave breaking over the reef flat [7]. The relative contributions of wave and tides on cross-reef exchanges are determined by both reef structure (size, submersion, roughness) and local hydrodynamic conditions [10, 17]. Thus, the understanding of reef hydrodynamics is of primary importance for estimating the cross-reef fluxes which affect the whole lagoon circulation and, consequently, the water renewal and biodiversity of the reef-lagoon system. Moreover, coral reefs themselves appear as fascinating examples of coupled bio-hydrodynamic feedback systems permanently evolving toward a dynamic equilibrium: the coral growth affects the hydrodynamics (turbulence, mixing, fluxes, nutrient transport and deposit) which partly controls the coral reef living conditions.

The knowledge of the cross-reef water fluxes and their impact on lagoon dynamics as well as their implementation in coastal circulation models are key issues for the estimation of flushing time and prediction of biological resilience of the complete reef/lagoon system. However, due to the toughness of measuring and analysing hydrodynamics at such a wide range of spatio-temporal scales, the cross-reef effects on the lagoon dynamics remain difficult to analyse and model at the whole lagoon scale. Tidally modulated wave-driven flows above reef barriers have been the main focus of several field studies [8, 10, 21, 7]. They all show that the tidal evolution of water depth over the reef top has a strong influence on the wave-induced phenomena. Analytical models proposed by Hearn [10] or Symonds et al. [8] relate the wave-driven current above the reef to the depth at the breakpoint and the water depth over the reef flat. Once these relationships have been calibrated thanks to a site-dependent empirical coefficient, a correct agreement has been found with field measurements in moderate wave conditions [7]. It should be noted that the calculation of the depth at the breaking point requires both a good knowledge of the external bathymetry and the definition of a breaking criteria. Contributions of cross-reef fluxes versus passages flows have been compared in micro-tidal conditions [22, 13], but few studies have evaluated the flows through reef passages and across the reef in macro or meso-tidal lagoons where the tidally induced flows across the reef are likely to be greater [23]. Studies of water exchanges between meso-tidal lagoons and open sea often focus on flows through reef passages only, as for the Noumea lagoon in New Caledonia [24, 25].

The long-term aim of our study is the characterization of the ecological and hydrodynamical functioning of a channel lagoon (see section 2.1 for details). The Ouano lagoon (New Caledonia) has been selected as a representative channel reef-lagoon system both exposed to Southern and Pacific Oceans swells and meso-tides range. As demonstrated by our field results (see section 3.3), tides and incoming offshore waves are the main forcings of the lagoon circulation and exchanges with the open ocean (through reef openings and above reef) and adjacent lagoon systems (through passages). The main objective of the present work is to improve, by the use of both field experiments and numerical modeling, our understanding of the role played by the cross-reef hydrodynamics on the circulation pattern at the whole lagoon scale. The field instrumentation is designed to analyse the exchanges processes while the numerical model is used as a tool for data interpretation and physical processes characterization. The latter does not aim, at this step, to provide a realistic description of the lagoon dynamics or to be used as an operational model. The adopted approach is, (i), to use a set of field measurements at strategic locations, i.e. at the lagoon boundaries, to analyse the response of ocean-lagoon exchanges to the variations of external forcings, (ii), to propose a simple parameterization of cross-reef fluxes based on the experimental results and, (iii), to implement and to test the relevancy of such parameterization in a coastal lagoon circulation numerical model used here as a numerical experimental tool in order to analyse the possible cross-reef flows influence on a narrow lagoon dynamics for which these effects are expected to have a major influence.

This article first describes the study area, the experimental strategy and the model. The field measurements section presents the main in-situ features of lagoon hydrodynamics. The next section is dedicated to the numerical study of the impact of cross-reef fluxes on the lagoon dynamics. Finally, the last section gives some conclusions and prospects, in particular concerning further numerical developments and combined hydro-biological research works.

2 Field site and methods

2.1 Field site

The Ouano lagoon, southwest New Caledonia, is used as an example of a coastal weakly-anthropized narrow lagoon (Fig. 1). Aiming to analyse the combined influence of waves and tides on the lagoon circulation, this site was specifically selected because it is both exposed to south pacific long swell waves and to meso-tidal fluctuations.

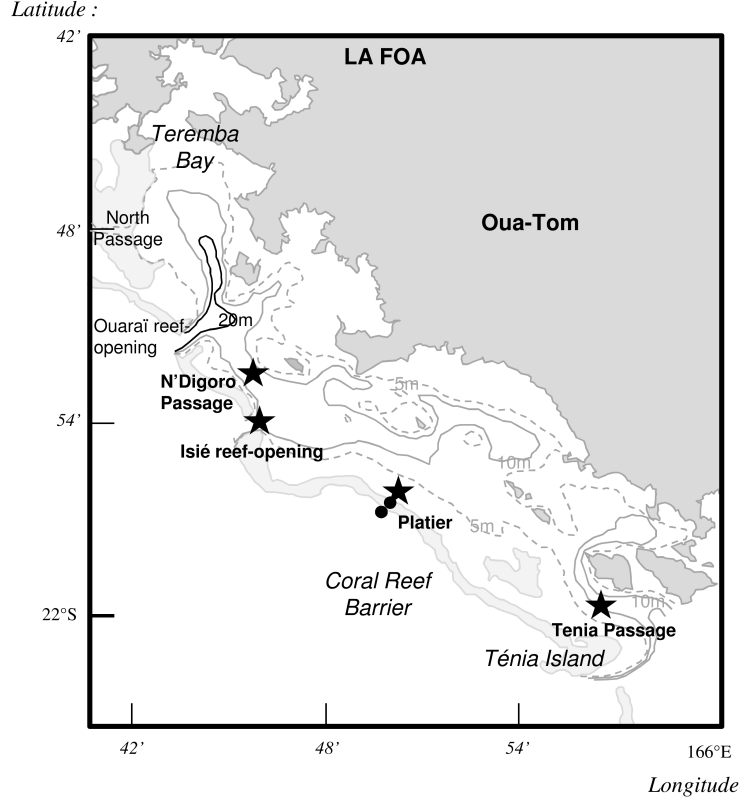


Fig. 1 Top view of the Ouano lagoon system. Black dots and stars represent pressure sensors and velocity profilers locations, respectively.

The Ouano lagoon can be classified as a channel lagoon, as often observed in high islands of tropical oceans. In their conceptual form, the channel lagoons are characterized by a shallow and narrow geometry and are free of any significant bathymetric sill observed in bowl-shaped lagoons. This particular structure promotes strong flushing currents forced by tides and waves and, consequently, rather short water renewal time.

The Ouano lagoon is approximately 30 km long, 10 km wide and 10 m deep. Its volume is about $2.1 \cdot 10^9 \text{ m}^3$. The lagoon is separated from the Pacific Ocean by a coral reef which is about 33 km long and varies from 6.3 to 10.1 km in width and it is directly opened to ocean through two reef-openings. The southern is about 1 km wide and 10-20m deep (the *Isié* reef opening) while the northern is the deepest, down to -60m and 1.5 km wide. The lagoon is connected to northern and southern lagoons by two passages, one toward north and one toward south. The northern passage is about 5m deep while the south passage (the *Ténia* passage) is 10-15 m deep. At high tide, the coral reef barrier is fully submerged, whereas at low tide it can be partly emerged depending on tide and wave conditions.

To facilitate the field experiment, the field study area is restricted to a nearly complete lagoon system extending from the *Ténia* passage to the N'Digoro passage where the lagoon topography is rather simple with a limited number of openings (see Fig. 1). In this part of the Ouano lagoon, the reef barrier is 25 km long and only opened at the *Isié* reef opening. The total volume of the considered portion of the lagoon is then $1.3 \cdot 10^9 \text{ m}^3$.

2.2 Field experiment and methods

The aim of the instrumental deployment is to characterize the exchange processes between the Ouano lagoon and the open ocean through reef openings (*Isié* site), above the barrier (*Platier* site) and with adjacent lagoons through *Ténia* and N'Digoro passages, with a specific focus on the response to offshore wave forcing. The instruments are thus located at the boundaries of the restricted lagoon system extending from the *Ténia* passage to the N'Digoro passage (see Fig. 1). The field measurement campaign has been

carried out during the 2013 winter austral season from August 28 to December 4. Four current profilers (ADCP) were deployed during the experiment to provide data on temporal variability of current velocity and direction along vertical profiles. Three Acoustic Doppler Current Profilers (ADCP) were deployed (Fig. 1) in lagoon passages (N'Digoro and Ténia) and reef-opening (Isié). An additional profiler was deployed in shallower area at the onshore end of the reef flat (the *Platier* site) to measure the cross-reef exchanges between lagoon and ocean. Mooring depths and profilers parameters are summarized in table 9. Note that instruments are mounted in heavily weighted cages or support brackets resting on the ground. Each presented mooring depth corresponds to the height of the water column above the sensor averaged over the whole acquisition duration.

Wave dynamics across the reef barrier was measured thanks to two high-frequency autonomous pressure sensors (OSSI Wave Gauge) laid on the bottom (with additional weight). The first one, called the *Ocean* site, is located over the reef outside slope ($14.40 \pm 0.015\text{m}$ deep) while the second one, called the *Reef* site, is at the landward end of the reef ($1.13 \pm 0.015\text{m}$ deep). Linear theory is used to estimate free surface oscillations from pressure measurements at the bottom. The mean water level, η , is obtained from a 30-min averaging of the 5Hz signal. The incoming significant wave height H_s is computed from the bottom pressure measurements at the reef outside *Ocean* sensor. Note that wave direction can not be extracted from such single pressure measurement point. Offshore wave data, including H_s^{ww3} significant wave height, T_p^{ww3} the peak period and θ^{ww3} the mean wave direction were computed by WAVEWATCH III model version 4.04. For a complete description of offshore wave computations, the reader is referred to the IFREMER wave hindcast runs for the IOWAGA project ([26]). Meteorological data (wind, pressure and humidity) during the field experiments were provided by the Tontuta airport station located over the same part of New Caledonia coastline, about 25km southwest of the Ouano lagoon and 16m over ground, and thus considered to be representative of the main wind features over the studied site.

Due to the overturning of the Platier ADCP during energetic conditions, most of data presented here is restricted to the August 28 - November 6 period. Such duration remains sufficient to cover a wide range of wave and tidal conditions.

All measured currents are projected along the main flow axis (see Section 3.2) to obtain the *main* and *transverse* components called U and V respectively. For the sake of simplicity, currents are considered positive (negative) when entering (leaving) the lagoon. The related flux is defined as the product of the depth-averaged main velocity by the local mean water depth: $Q = UH$.

2.3 Numerical modelisation

2.3.1 The numerical tool : ROMS

The Ouano numerical modelling was performed with the ROMS (Regional Ocean Modeling System) three dimensional dynamic model [27]. This free-surface ocean model solves the Navier-Stokes equations associated with the classical Boussinesq and hydrostatic pressure hypotheses. The sub-grid scale parameterization uses the concept of turbulent viscosity and diffusivity. Here, the Smagorinsky parameterization has been used for horizontal mixing and a LMD-KPP closure parameterization [28] has been used for the vertical mixing.

Similar to other oceanic models (as POM, [29]), ROMS is a split-explicit model, where short time steps are used to advance the surface elevation and barotropic momentum according to the CFL stability condition and the full 3D equations are solved separately with higher time steps. Furthermore, ROMS employs a special 2D averaging procedure for barotropic mode, which satisfies the 3D continuity equation exactly in a finite-volume/finite-time-step sense and guarantees exact volume-conservation and consistency preservation for the 3D.

2.3.2 Bathymetry and discretization

Due to the absence of refined local bathymetric data for the studied lagoon, the bathymetry used here is an interpolation of the 131 000 depth soundings performed by the Service Hydrographique et Océanographique de la Marine (SHOM/MOP) over the whole New Caledonia lagoon between 1960 and 1985. The measurements horizontal resolution is not constant over the lagoon, varying for a few tens of meters to several hundreds of meters in the hard-to-access regions. As the initial aim of the SHOM's soundings was to provide navigation maps, some areas of the Ouano lagoon are poorly discretized in particular in shallow near-reef regions, in the La Foa bay or in the northern part of the lagoon. The interpolation of the SHOM's soundings, accounting for the data density, has been carried out using the UNIRAS software library from

Advanced Visual Systems. Whereas finer and complete bathymetric measurements of the Ouano lagoon should be performed in a more operational context of numerical modeling, the present bathymetry appears sufficiently accurate to run our schematic model for process studies.

Spatial discretization is based on the finite differences using an Arakawa-C grid combined with topography-following coordinate in the vertical direction (σ coordinate) discretized into 10 vertical levels with surface refinement. The study area was discretized horizontally using a uniform mesh of 108×192 cells of about 250 m resolution. This is generally of the same order of precision than the soundings resolution. The reef barrier extends thus over 2 to 4 meshes. A tricky issue is the characterization of very shallow regions, in particular the shoreline and the reef barrier itself. Shoreline and coral reef limits are determined thanks to the New Caledonia government lagoon map [30]. A flat bathymetry is imposed over the reef barrier, which do not represent the very complex and rough structure of the real reef. The vertical elevation of the reef flat is imposed from in-situ observations: the reef is just emerged during low waters of extreme spring tide. The topography between the reef center and the lagoon has been slightly smoothed to represent the transitional region from the reef flat to the lagoon sand bottom which is poorly documented by the bathymetric soundings.

Passage cross-sectional areas are estimated from the present bathymetry to compute exchanged volumes. For the N'Digoro, Platier, Isié and Ténia sites, they are about $2.09.10^4$, $3.53.10^4$, $1.15.10^4$ and $2.85.10^4$ m², respectively.

The extension of outside open ocean domain is set to be large enough to avoid the influence of boundary conditions (see next section). A threshold for maximal depth is imposed at 50m in order to minimize time steps constraints by depth.

The time step is constrained by the CFL condition. For the no-wave reference case, it is mainly determined by the tidal wave celerity in the open ocean domain. When wave-induced fluxes are imposed on the reef barrier (see below), the time step is first limited by the velocity over the reef which can sometimes increase above $1m.s^{-1}$ in particular during model spinup. In the present study, the time step varies from 4s for the *Reference* and *Wall* simulation to 2s for the 1.2m wave case simulation. Further numerical work will certainly allow to optimize these values.

2.3.3 Boundary conditions and external forcings

The model domain is a rectangular area of about 26×48 km, including the Ouano lagoon. It is oriented following the main coastline direction. Hence, the north-east boundary is the land whereas the south-west boundary is the offshore boundary, about 25 km far from the coast. The lateral boundaries are located southward the Ténia Passage and northward the north passage.

Tides at open boundaries are calculated with ADCIRC (ADvanced CIRCulation model, [31]) over the New Caledonia region on a fine unstructured grid (from 500 to 25 m in shallow water). Only one tide component is imposed here, the M2 wave. As suggested by Bourret et al. [32] who studied the influence of open boundary conditions on coastal circulation, tidal forcing was specified at the offshore (south-west) boundary with a Neuman condition while water level at lateral boundaries (north-west and south-east) has been determined with a Chapman condition [33]. Velocities are deduced from a Characteristic Method [34] in the offshore boundary and with a Flather condition [35] in the lateral boundaries.

As the present study focuses on the role played by wave-induced fluxes on the lagoon dynamics, the effects of stratification, river inputs and atmospheric forcings (thermal fluxes, rainfall, wind, pressure gradients) have been neglected in the present numerical simulations.

At the free surface, the wind stress is neglected and so is the vertical shear of horizontal velocity. At the bottom, the bottom stress is classically prescribed with a quadratic law formulation controlled by C_d the bottom drag coefficient and \vec{U}_b the bottom velocity near the bottom, i.e at the center of the bottom mesh [36] :

$$\vec{\tau}_b = \rho C_d \|\vec{U}_b\| \cdot \vec{U}_b \quad (1)$$

where $\vec{\tau}_b$ is the bed shear stress.

The drag coefficient is kept constant as the vertical resolution is not sufficient to resolve the bottom boundary layer [29]. However, in order to characterize the difference of bottom structure and characteristic roughness heights, two values are used. For the sandy regions of the domain, the classical value $C_d = 0.0025$ is used [29]. Over the reef, strong roughness effects of coral reef on boundary layer and friction processes are expected [20]. A set of calibration tests on the M2 tide component have been performed to determine the most relevant reef bottom drag coefficient (see section 4.1.1). The value $C_d = 0.01$ is used, which is very close to the one measured on the field by Reidenbach et al. [12].

2.3.4 Initial conditions for the hydrodynamic simulation

Velocities and water level were set to zero at the beginning of the simulation. About 4 days are required to spinup the model and cancel the influence of initial conditions. Most of simulation lasted for 40 days and only the 20 last days have been used for numerical data analysis.

2.3.5 The test cases

As explained in introduction, the simulations are used as a process-study tool to understand the impact of cross-reef fluxes on channel lagoon dynamics and not to provide an operational model of the Ouano lagoon. Our main aim being to discriminate the role played by tides and variable wave forcing on the exchange fluxes through passages and openings and over the reef barrier, all other external inputs (wind, rainfall, etc) are neglected (see section 2.3.3). By the same token, only one tide component is considered (M2 tide).

In order to point out the influence of wave-induced cross-reef water fluxes, three schematic types of simulations have been carried out:

- The *Reference* simulation. In this simulation, the reef is simply considered as flat portion of the bathymetry, raised to a constant given elevation (see section 2.3.2) and characterized by an increased drag coefficient (see section 2.3.3). No wave forcing is considered.
- A *Wall* simulation. In this simulation, the coral barrier has been considered as an impervious barrier inhibiting any flow above the barrier. This case is compared to the *Reference* simulation in order to test the effect of water exchanges above the reef on the overall lagoon dynamics in the absence of wave.
- The *Wave* simulations. They are based on the *Reference* simulation for which wave are supposed to break on the reef. Wave height are kept constant during each simulation and varies from simulation to another from 0.1m to 1.2m. Water fluxes induced by wave breaking on the reef are parameterized in the model using the empirical law determined from in-situ measurements (see section 3.4)

2.3.6 Numerical product: Origin and age of lagoon water

Several research works have been dedicated to the numerical characterization of lagoon waters renewal time, see e.g. [4, 37–41]. In our case, we use the age as a relevant quantity to estimate the dynamics and history of lagoon waters defined by De Brye et al. [40] as "the age of a water parcel located at point x and time t is the time elapsed since the water parcel entered the domain of interest". The age is calculated inside our computational domain divided in three areas:

- The domain of interest : the lagoon area extending from the mainland shoreline to the coral reef barrier and from the Ténia passage to the North passage
- The ocean area, out of the reef barrier
- The neighboring lagoons (south-east and north-west)

The age is a Lagrangian concept; however, as suggested by De Brye et al [40], it can be estimated from an Eulerian approach (see [40], eqs. 1-3). Using this method, both original location and age of lagoon water can be determined. Note that the lagoon renewal time water could be estimated as the moment when the whole water mass contained inside the lagoon has an external origin. However, in the present paper, a focus is made on the above-mentioned quantities (age and origin) to characterize the lagoon renewal dynamics (see sec. 4.3). As constant forcing (waves and tides) is applied for each simulation, the computation of age and origin of lagoon water shows an asymptotic behaviour. The simulations are thus set to last until all lagoon water is renewed (until few hundred days for some simulation cases) and asymptotic values reached by both age and origin are extracted and analysed, see section 4.3.

3 Field measurements results

The overall objective of the field part of our work is to provide a set of in-situ measurements to identify the forcing agents of the lagoon dynamics and to understand their coupled influences. We first focus on the wave transformation over the reef and then describe the main trends observed at the measurement sites in terms of levels and currents.

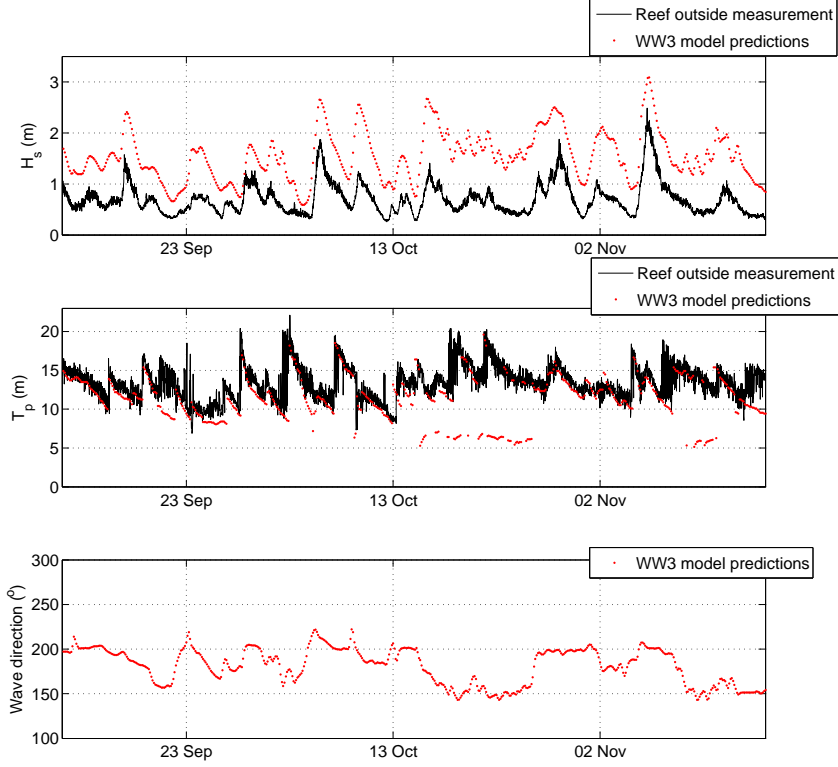


Fig. 2 Wave features during the studied period. Top plot: comparison between significant wave heights measured at the Ocean sensor (black line) and provided by WW3 model (red dots). Middle plot: comparison between significant wave heights measured at the Ocean sensor (black line) and provided by WW3 model (red dots). Bottom plot: Wave direction provided by WW3 model.

3.1 Wave dynamics

The characterization of wave propagation and transformation above the reef barrier is of great importance to understand the energy and mass transfers from open ocean to lagoon. Wave measurements performed at the Ocean site (see Fig. 2) are typical of long swell waves, with significant wave height ranging from 0.4 to 2.3m and most of measured peak period ranging from 10 to 20s. Several energetic swell events are observed during which H_s remains higher than 1.2m for several hours or days. In-situ wave measurements are compared with WW3 numerical model estimations [26], see Fig. 2. The significant wave heights comparison (top plot) shows that, even if a strong correlation exists (correlation coefficient 85%), the numerical estimations systematically overestimate the value measured at the reef outside slope: the ratio numerical and measured data ranges from 1.1 to 5.1 with a mean value around 2.4. As observed when comparing the overestimation factor with the predicted wave direction, this trend may partly be attributed to refraction effects: the best agreement between model and measurements is obtained for normally incident wave (the mean reef angle is nearly 210°). Further field measurements have to be performed to better understand such discrepancy between model and observations and characterize the wave transformation on the reef outside slope including bottom friction and reflection processes.

A representative 6-day time period from Oct. 3 to Oct. 9 is selected to emphasize the wave transformation above the reef. Figure 3 compares the time evolution of the spectral density of energy for the Reef and Ocean sensors, according to the corresponding 30-min averaged water level.

Wave dynamics at the outside reef (Ocean site, right plot) is focused in the gravity band. Most of the energy ($> 95\%$) is typical of long swell waves, with corresponding wave periods between 8 and 22s, and very weak energy for high frequencies. A significant swell event, with significant wave height two times higher than the average height over the whole experiment, is observed to hit the reef barrier on Oct. 5 and slowly decay on the following days. No tide dependency is observed. The Reef sensor (Fig. 3, middle plot) reveals radically different features. The wave dynamics over the reef is strongly controlled by the

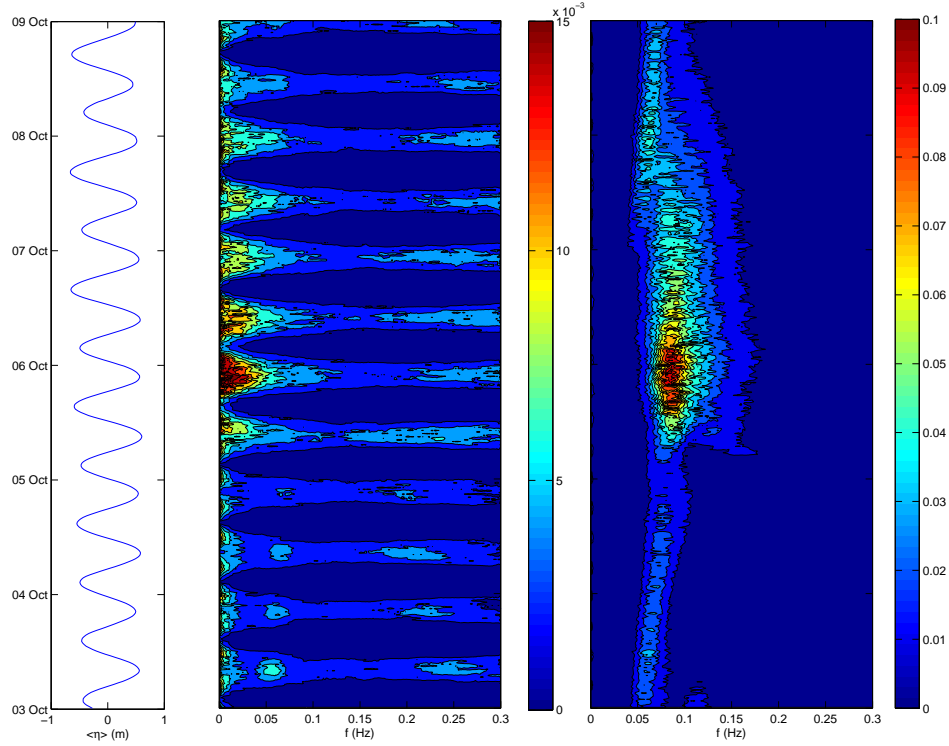


Fig. 3 Wave transformation above the reef barrier. Left plot: 30-min averaged water level. Middle plot: Spectral Energy Density at the Reef sensor in $m^2 \cdot Hz^{-1}$. Right plot: Spectral Energy Density at the Ocean sensor in $m^2 \cdot Hz^{-1}$. Note that, as strong energy dissipation occurs during wave breaking and propagation over the reef, the spectral densities of energy at the Ocean and Reef sensors show quite different ranges (nearly one order of magnitude). Different scales are thus used for the sake of visual commodity.

tide level: the higher the mean level is, the stronger the wave energy is. A large part of the wave energy (around 45% for $H_s < 0.5m$ and up to 85 % at the swell peak $H_s = 1.7m$) is here distributed in the band $0 < f < 0.05Hz$. This highlights the strong energy transfer toward long infragravity waves during swell wave breaking above the reef as described by Pomeroy et al. [14]. Some short wave energy is also observed in the band $0.15 < f < 0.3Hz$. Estimation of wind-generated waves using the typical lagoon length-scale [42] and wind measurements from the Tontouta airport station 25km southward from the Ouano lagoon) provides wind wave frequency from 0.4 to 0.8Hz, which is much higher than the frequencies observed here. The most probable source for the observed free surface motion in the $0.15 < f < 0.3Hz$ band is the generation of higher harmonics (i.e. secondary waves) as the incoming gravity swell propagates over the reef [43].

It is also noted that small and very long swell waves observed the first two days at the Ocean sensor are more prone to propagate above the reef than shorter but much higher waves observed from Oct. 5 to Oct. 8. For such large swell events, the strong energy peak in the swell gravity band at the reef outside is nearly absent on the reef. As explained above, the incoming energy has been mostly transferred toward low frequency infragravity motions in that case. Additional field measurements are necessary to better understand the wave transformation above the reef barrier in the continuity of the works recently engaged [7, 14]. This will require a high resolution measurements both in time and horizontal direction along the selected cross-reef transect from the shoaling zone above the outside slope up to the inner part of the reef in order to, (i), discriminate the roles played of reflection, dissipation and harmonic transfers and, (ii), improve the definition of breaking threshold which is of primary importance for the estimation of wave-induced fluxes (see section 3.4).

	Main angle	U_{max}	U_{min}	\bar{U}	$ \bar{U} $	V_{max}	V_{min}	\bar{V}	$ \bar{V} $	$\frac{ \bar{V} }{ \bar{U} }$
N'Digoro	327	0.29	-0.63	-0.11	0.14	0.11	-0.13	0.001	0.014	0.09
Ténia	123	0.39	-0.44	-0.03	0.1	0.13	-0.2	0.003	0.015	0.15
Platier	10	0.44	0.07	-0.07	0.07	0.19	-0.19	0.006	0.036	0.5
Isié	265	0.7	-0.6	-0.06	0.21	0.23	-0.14	0.02	0.04	0.19

Table 1 Depth-averaged velocity statistics for the four measurement sites ($m.s^{-1}$): maximal value, minimal value, mean value and mean value of the magnitude for both main and transverse components and ratio between the mean values of magnitude.

3.2 Levels and currents

As expected for meso-tidal lagoon, free surface motions are characterized by a semi-diurnal fluctuation with an amplitude varying from 0.2 to 0.7m during the neap and spring tides and the water level is generally the same at all moorings location.

Velocity profiles showed a quite small vertical variability (less than $5 cm.s^{-1}$ on average from up to bottom velocity), thus the depth-averaged values are considered as representative. Statistics for depth-averaged velocities are shown in table 1. Let us first focus on the hydrodynamics within passages and openings. The depth-averaged currents in the passages and reef opening have been observed to follow a main direction, either in- or outward. As illustrated by the dispersal ellipses (Figure 4), these main directions are 123° , 327° and 265° for the Ténia, N'Digoro and Isié sites, respectively. The dominance of the main component is demonstrated by the ratio between temporally averaged values of the magnitudes of both velocity components $|\bar{V}|/|\bar{U}|$ (see table 1, last column). This ratio ranges between 0.09 and 0.19, which indicates a strong dominance of the main component for the Ténia, N'Digoro and Isié sites. The usual tide-related oscillations with in- and out-flows are observed at each mooring. As highlighted by the dispersal ellipses asymmetry and the negative mean values of the main component, the tide-averaged dominant trend for Ténia, Isié and N'Digoro sites is an outflow leaving the lagoon. However, in- and outflows are rather well balanced for Isié and Ténia, inducing weak (outflow) time-averaged velocities at these locations (around $5cm.s^{-1}$) whereas maximal outward flow at the N'Digoro site is about two times greater than the inward current and the mean velocity is above $10cm.s^{-1}$. The Isié mooring shows the greater velocities both in instantaneous and averaged values, reaching up to 0.7 and $0.21 m.s^{-1}$, respectively. The mean current values are lower in N'Digoro ($|\bar{U}| = 0.14 m.s^{-1}$) and Ténia ($|\bar{U}| = 0.07 m.s^{-1}$).

The current at the Platier site is also along main direction, but the transverse component is here generally more important than in other sites ($|\bar{V}|/|\bar{U}| = 0.5$). The mean value for main component is positive revealing a dominant inflow tendency (93% of the sequence). The depth-averaged velocity mainly ranged from about 0 to $0.2 m.s^{-1}$ (89% of the collected data). The currents at the Platier site appear to be strongly related to the external wave forcing. Rare but high instantaneous entering current larger than $0.3 m.s^{-1}$ are observed during large wave episodes ($H_s > 1.5m$, which occurs about 1.5 % of the time) whereas weak velocity fluctuations around zero ($|\bar{U}| < 0.05m.s^{-1}$) are observed during calm periods ($H_s < 0.5m$ with occurrence frequency about 26 %).

3.3 Dynamics of the reef-lagoon system

Usually, water circulation within lagoons is mainly controlled by fluxes through passages, reef opening and above the immersed part of the coral barrier in response to external forcings, such as tides, waves, wind, etc [7,17,44,45]. In the considered lagoon, the main driver is the tide, as demonstrated for instance by the dispersal ellipses (Fig. 4) or the periodic oscillations of velocities observed in Fig. 5. In order to identify the relative importance of the other forcings on the lagoon circulation, velocity data have been bandpass filtered to remove the influence of the M2 and S2 tidal components. Correlations between tidal-filtered depth-averaged velocity and external parameters are shown in table 2. We only consider here the effects of incoming wave height H_s and peak period T_p , the wave angle θ_{ww3} and the wind magnitude W and direction θ_w . Rainfalls have been sparse on the studied period: only ten rainy days with cumulated rainfall on the total campaign about 70mm and a maximal daily value about 12mm. The effect of rain is thus neglected. Wind can also be expected to affect the lagoon-reef hydrodynamics, depending on the local wind intensity. Wind measurements performed at the nearby Tontouta airport station revealed the typical wind pattern observed during the experiments. Trade winds are modulated by thermal breeze and guided by

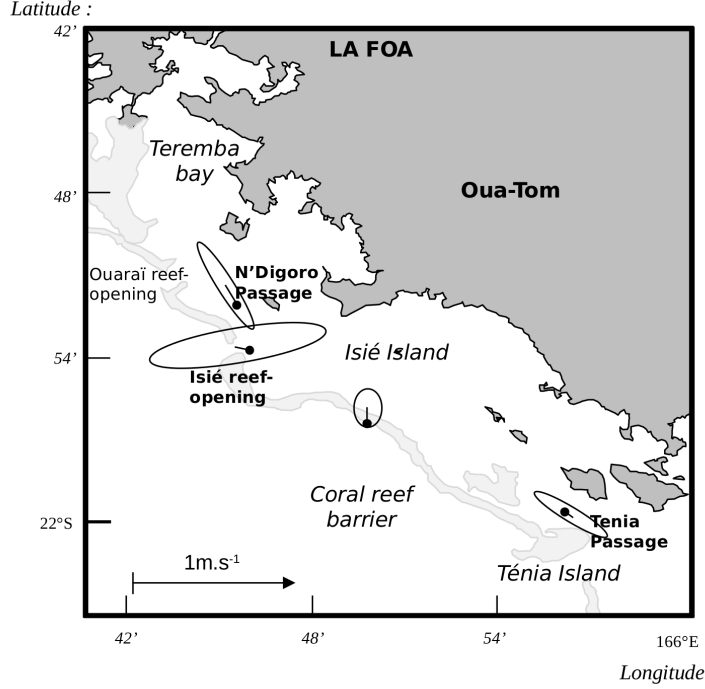


Fig. 4 Dispersal ellipses for the four mooring sites (black dots represent the zero values and scale is given by the 1m.s^{-1} arrow)

Designation	N'Digoro	Ténia	Platier	Isié
H_s	57	21	80	62
T_p	1	19	15	5
θ_{ww3}	7	50	30	4
W	21	30	1	19
θ_w	11	10	6	14

Table 2 Correlation (in %) between tidal-filtered velocity and external parameters at the four measurement sites. H_s and T_p are the significant wave height and wave period measured at the Ocean sensor, θ_{ww3} is the offshore wave direction computed by WW3 model and W and θ_w are the wind magnitude and direction measured at the Tontouta airport station

the mountainous topography of New Caledonia. Winds are minimal during nights (lower than 1 m/s) and nearly offshore (north-west to north). They increase during the day in the range 3-6 m/s while clockwise rotating and blowing from the east, then south and finally west before slowing down in late afternoon.

Table 2 shows that, apart from the tide effect, the incoming wave height is the principal external parameter affecting the lagoon dynamics in the studied period. As expected, this wave effect is overwhelming at the Platier site which is directly exposed to wave breaking induced water fluxes. The wave height is also well correlated ($> 50\%$) to the tidal-filtered dynamics at the N'Digoro and Isié sites. Other correlation coefficients are too small to infer a clear relationship between measured currents and other external parameters at Platier, N'Digoro and Isié sites. Velocity measurements at the Ténia site reveal a different behavior. It is not dominated by incoming wave height effect, but shows significant dependence to wave direction and, more weakly, to wind magnitude.

In our following attempt to parameterize wave effect on the lagoon dynamics, the choice is made to focus on the significant wave height which is, apart from the tide, the dominant external driver of the velocity in three out of four of the studied sites. The role of incoming wave direction is neglected, but it probably deserves detailed consideration in further improvement of the wave effect parameterization. The direct influence of the wind stress is also neglected, as it was in other lagoon studies [46, 47, 13] in similar wind conditions, but should also be considered as a secondary forcing in realistic numerical approach, which is not the aim of the present paper.

3.3.1 Tidal influence

The power spectrum estimated by periodogram analysis [48] of the amplitude of water level variation confirmed the prevalence of the semi-diurnal and diurnal tides M2, S2 and K1 which determine about 97% of the signal. The tidal range varied slightly inside the lagoon and there was a weak phase lag of few minutes (above the error [48] between each station in table 10).

The tide also determines, but in a lesser extent, the magnitude and direction of the currents in the lagoon. Spectral analysis [49,50] of the velocity measured at the moorings confirmed the dominant role played by semi-diurnal and diurnal components M2, S2 and K1 on the velocities and also revealed the presence of quarter-diurnal component (M4). Their influence is really dominant for the current through the Isié reef opening where the tidal current represents about 93% of the total current, but also in Ténia and N'Digoro passages where tidal currents about 66% of the magnitude of the observed instantaneous currents, see table 10. At the Platier site, this tidal influence is less visible.

Whereas the intermittent permeability of the barrier tends to cancel out the phase lag between the cross-reef current and the sea surface height, currents at moorings in the reef openings and passages are driven by the difference in water level inside and outside the lagoon. At Isié, N'Digoro and Ténia sites, depth-averaged velocities were seen to be nearly in quadrature with the water level: The phase lag ranged from 2.5 to 3.5 hours for the M2 tide (table 10) with an off-shore to in-shore progression, for the M2 component wave of tidal flow. The flood starts at reef opening, then is observed in N'Digoro passage and finally, few minutes later, at Ténia. Hence, the induced tide inside the lagoon was, a forced almost standing wave, with few minutes of phase lag.

The lagoon is observed to fill (resp. empty) with the flood (ebb) tidal cycle through both passages and reef opening. About 30 % of the tidal fluxes transits, each tide, through Isié reef opening; 20 % through N'Digoro passage and above the reef, but only 2 % through Ténia Passage. These proportions slightly varies with tide amplitude. One notes that, in comparison with neap tides, the part of the exchanged flux at N'Digoro with respect to the total flux increases during spring tides of about 20% while its counterparts in Platier and Isié sites decrease of about 5%.

The lagoon volume ranged from about 1.2 to $1.5 \cdot 10^9 m^3$ and from 1.3 to $1.4 \cdot 10^9 m^3$ during spring and neap tides, respectively. Hence, using the tidal prism method [51], the tidal-prism flushing time can be estimated at 2 days during spring tide but only 8 days during neap tide.

3.3.2 Tidally-averaged dynamics

Depth-averaged velocity measurements can be averaged in time over the successive tidal cycles to remove the influence of tide oscillations on the water fluxes. Applied on the Platier profiler data, this tide-averaging reveals the relationship between cross-reef current and wave forcing (with a correlation of 0.81, see section 3.4 for further analysis of cross-reef fluxes). Averaged on the campaign, the mean cross-reef current is estimated to $0.07 m.s^{-1}$. As it takes place all over the reef flat, this flux plays a significant role in the total exchanges between lagoon and open sea. Indeed, integrated all along the reef this inflow represents about 40% of the lagoon volume each day. Such input is compensated by strong exchanged volume in Isié reef opening (about $0.5 \cdot 10^8 m^3$ a day) and passages ($2 \cdot 10^8 m^3$ and $0.8 \cdot 10^8 m^3$ a day through N'Digoro and Ténia passages, respectively). The wave-driven cross-reef exchanges are enhanced during strong wave activity, as the one observed the October, 5-6 (see next section). During this event, the tidally-averaged velocity above the reef can reach about $0.15 m.s^{-1}$, and induces an exchange of about $11 \cdot 10^8 m^3$ a day, which is about 85% of the volume of the lagoon in one day.

3.3.3 Lagoon circulation: a coupled tide-surface-wave dynamics

As explained above, tides and waves are by far the dominant drivers of the studied lagoon circulation. To highlight their coupled influence, two 24-h measurements periods are presented in Fig. 5. The first period (left plots) lasts from Oct. 3, 23:00 to Oct. 4 23:00 while the second period (right plots) lasts from Oct. 5, 12:00 to Oct. 6 12:00. Both show a nearly similar tidal range (0.6m) but quite different significant wave heights (0.3m and 1.2m). Top and bottom plots represent the flux at the four measurements sites and the 30-min averaged sea surface level measured at the Ocean sensor, respectively. The red dots in bottom plots indicate if the breaking criterion threshold is reached.

During moderate wave conditions (Fig. 5, left plots), one first notes that the fluxes is maximal in the Isié reef opening, which is directly connected to the open ocean, whereas the smallest flux is observed for the Platier site which measures the cross-reef current. The temporal evolution of flux highlights both phase

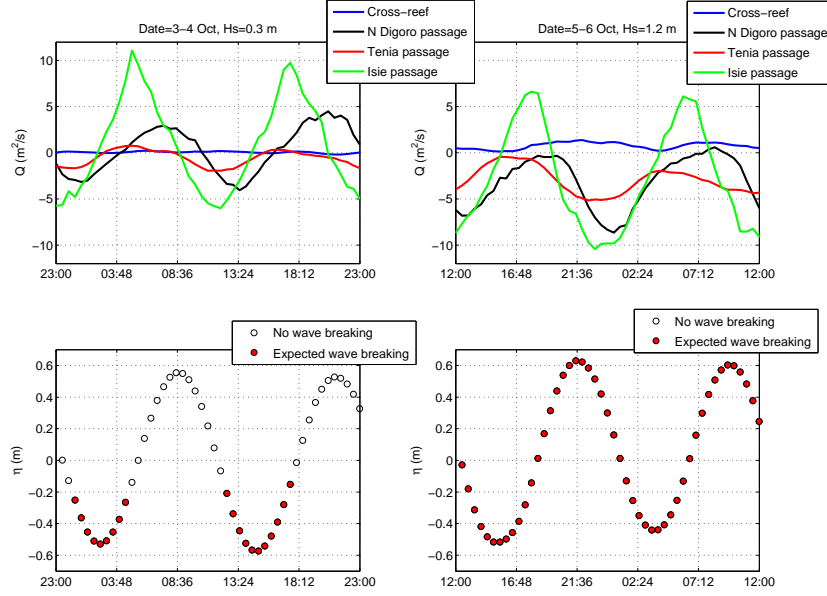


Fig. 5 Comparison between calm and strong wave events on the left and right plots, respectively. Top plots depict the fluxes calculated from depth-averaged velocity at each site while bottom plots shows the mean water level fluctuations with estimation of wave-breaking occurrence (red dots) using ad-hoc breaking criterion.

shifts and differences in tidal dynamics at the studied sites. At the Isié passage, the flow reversals around the tide apices are much less smooth than at the other sites which show the typical sin-like evolution of tidal flows. An important observation is that, for these calm wave conditions, fluxes at the N'Digoro passage are nearly symmetric (similar flood and ebb magnitudes), while the Isié passage shows a more pronounced inflow. This trend is even clearer at the Ténia site for which small inflows are observed only during the first part of the rising tide.

When the reef barrier is exposed to stronger wave energy (Fig. 5, right plot), the overall scheme appears to be similar, in terms of phases and oscillation magnitude, but a dramatic shift is observed toward an emptying tendency through each passage. Through N'Digoro and Ténia passages, the flux is systematically directed toward the open ocean, even at rising tide, while at the Isié reef opening, the outflow becomes clearly dominant. These observations should be related to the significant cross-reef inflow occurring above the barrier. The offshore wave energy is here sufficient to produce a permanent wave breaking all along the tidal oscillations. The related induced current drives water into the lagoon all over the reef barrier which finally forces a nearly generalized outflow toward the passages.

This again demonstrates the major role played by the wave-induced current into the water masses dynamics at the lagoon scale.

3.4 Cross-reef Fluxes

In most of existing research works on cross-reef fluxes [8,7,10], the flux calculation generally requires the description of wave transformation during shoaling and breaking processes and, in particular, to compute the wave height at the breaking point. Such quantity can be obtained using linear shoaling theory and breaking criterion [7] but this would require a high-resolution (meter scale) knowledge of the bathymetry in cross-reef direction and can hardly be generalized all along the reef. A more global criterion is, then, expected. The choice is made here to only relate the flux to easily available external parameters. Preliminary analysis demonstrated that the depth-averaged current is primarily dependent on the water depth above the reef and on the incoming significant wave height which were chosen to be the main governing external parameters.

A simple breaking criterion based on the ratio between the incident significant wave height H_s and the depth over the reef flat H is defined: when $H_s/H > 0.7$ waves are expected to break over the reef flat, whereas when $H_s/H < 0.7$, no wave breaking is expected. The difference between breaking and non-

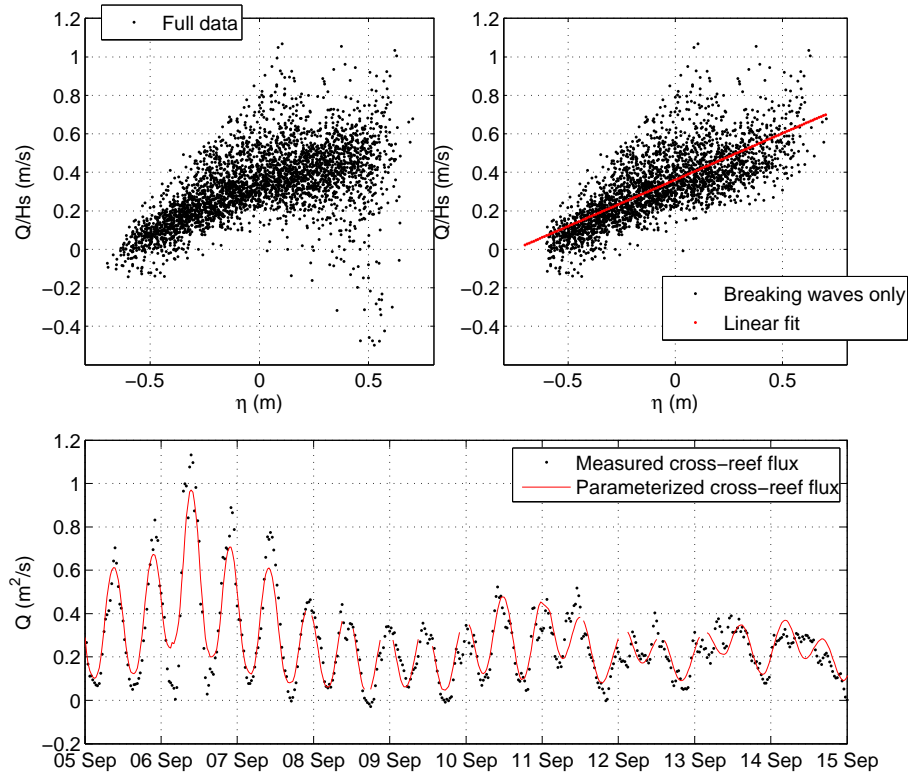


Fig. 6 Cross-reef flux parameterization. The top plots depicts the evolution of the cross-reef flux divided by the significant wave height Q/H_s with respect to the mean water level η . Full data is shown in top left plot while and the data where breaking is expected is show in top right plot. Bottom plot represents for a ten days time period the comparison of measured and calculated fluxes. The discontinuities in calculated flux correspond to the non-breaking periods. Positive velocity values correspond to inward current entering the lagoon.

breaking conditions clearly appears in the cross-reef current (Fig. 6, top plots). At high tide, in wave breaking conditions, the flux tends to be relatively high, whereas it is weak or negative (offshore directed) when no breaking occurred.

This selected breaking threshold uses the same value that the one proposed by Sysmonds et al [8] although their criterion is based on the wave height calculated at the breaking point, not the incoming significant wave height measured in 14.4m depth as in our case. In the continuity of field measurement works recently engaged [7–14], further high (horizontal) resolution experiments should be carried out to provide a finer characterization of wave transformation over the reef barrier. This requires to deploy high-frequency (wave-scale) instruments from the outside slope up to the lagoon to better understand shoaling, reflection, wave breaking, dissipation and harmonic transfers with a sufficient horizontal resolution to capture surf zone processes which shift with tides and incoming wave height. Additional measurements should include other locations along the reef to estimate the longshore variability of the observed processes. A particular attention must be paid on the estimation on breaking threshold, which has been here indirectly characterized using the dependence of the cross-reef current to the H_s/H ratio. Nonetheless, it should be emphasized that, in nature, the initiation of breaking for real wave spectrum over a complex bathymetry such as coral reef is not an sudden process, but rather a progressive transition from partial breaking to a fully saturated surf zone. The triggering for the activation of wave breaking induced flux in the model, described later on, should be a smooth process and not drastically sensitive to the exact value of breaking threshold. Note also that, aiming to implement the cross-reef flux parameterization in a lagoon circulation numerical model, a constant bed elevation value is used to average the very complex bathymetry of coral reef barrier.

	Amp ^{Model} (m)	Amp ^{Meas.} (m)	Phase ^{Model} (°)	Phase ^{Meas.} (°)	Corr. (%)	Error
N'Digoro	0.40	0.38	272	272	100	2.10^{-4}
Isie	0.40	0.36	272	272	100	6.10^{-4}
Reef	0.39	0.40	270	272	100	2.10^{-4}
Ténia	0.40	0.41	272	272	100	$0.5.10^{-4}$

Table 3 Comparison between model computations (*Reference* simulation) and measurements for the M2 tidal component of water level at each mooring. Tidal amplitude and phase, correlation and mean squared error.

	Magn ^{Model} (m/s)	Magn ^{Meas.} (m/s)	Phase ^{Model} (°)	Phase ^{Meas.} (°)	Corr. (%)	Error
N'Digoro	0.03	0.08	82	198	-44	$5.0.10^{-3}$
Isie	0.25	0.29	173	166	99	$1.5.10^{-3}$
Reef	0.07	0.03	64	269	-90	$4.4.10^{-3}$
Ténia	0.13	0.09	165	203	78	$3.2.10^{-3}$

Table 4 Comparison between model computations (*Reference* simulation) and measurements for the M2 tidal component of velocity at each mooring: Current amplitude and phase, correlation and mean squared error.

Under wave breaking conditions (Fig. 6, right plot), the flux tends to be a monotonically increasing function of the water level, η , and, even if a rather large variability is observed, a fitted linear relationship may be obtained:

$$Q = H_s(A * \eta + B) \quad (2)$$

where $A = 0.48s^{-1}$ and $B = 0.36m.s^{-1}$ are empirically-determined coefficients by least-square method.

As illustrated in figure 6 where measured and parametrized cross-reef velocity are presented during a period with both high (Sept. 5-8) and low (Sept. 12-14) energy events, this empirical linear relationship provide a quite fair representation of the order of magnitude and main tendencies of the cross-reef flux even if the extreme values at the tide reversals can be underestimated. The correlation coefficient between the linear fit and the measured data in breaking conditions is 0.74. As expected, such agreement is slightly lower than the correlation obtained, for instance, by Bonneton et al. [7] which account for more physical considerations in their cross-reef fluxes parameterization (see also [10,8]). However, their analysis requires the computation of internal parameters of the system, such as the wave setup or wave height at the breaking point, which necessitates a detailed knowledge of the bathymetry (see details above). The present results suggest that, in the context of lagoon circulation studies, our simple parameterization in terms of the sole external parameters provides a crude, but easy-to-use description of the cross-reef fluxes and that the benefits of using finer fluxes estimation may be offset by the need to measure the bathymetry at a sufficient resolution to compute wave features up to the breaking point using linear wave theory [7] and the cost of additional computations.

4 Impact of the cross-reef dynamics

4.1 Numerical developments

4.1.1 Validation of the Reference simulation

The *Reference* simulation is first used to validate our numerical choices, in particular in terms of boundary conditions and bottom drag coefficients (see section 2.3.3), and then to simulate the tidal propagation within the Ouano lagoon. Recalling that numerical simulations are used here as a tool for process study and not as a fully realistic model of the Ouano lagoon, the validation has to be regarded as a verification of the model ability to simulate the tidal dynamics. The analysis focuses on the M2 tide component, in terms of amplitude, phase and related currents at each of the selected moorings. Tables 3 and 4 show the comparison between numerical results for the *Reference* simulation and in-situ tidal components extracted over the whole field data (see section 3.3.1 and table 10) for the water level fluctuations and tidal-induced currents, respectively.

A very good agreement is noted for the tidal free surface dynamics (table 3) both in terms of amplitude, correlation and phase. The simulation of tidal currents shown in table 4 is less accurate and calls for comments. Currents magnitude and phase are rather well reproduced by the model at the Isié and Ténia

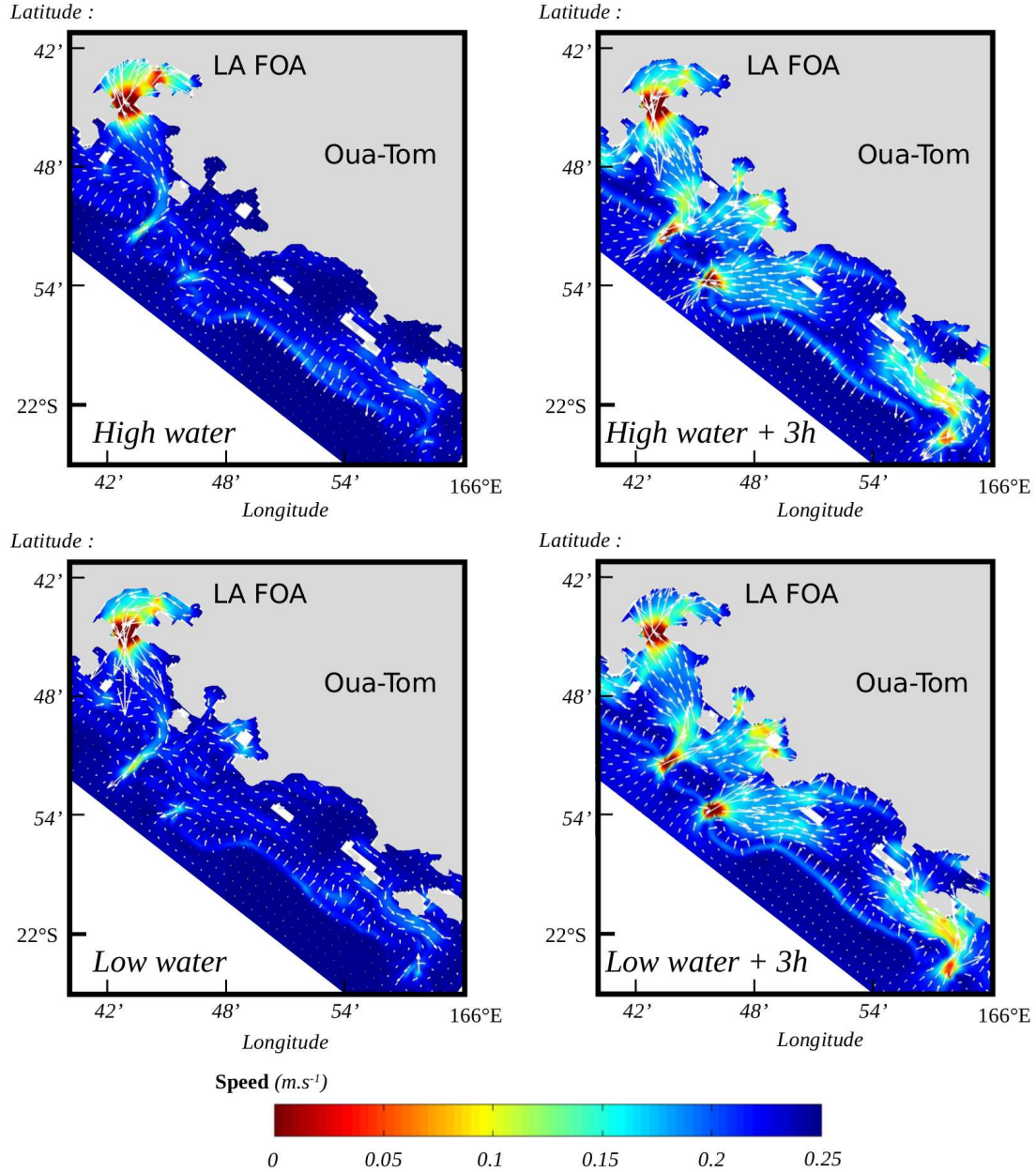


Fig. 7 Numerical lagoon dynamics for the *Reference* case, only tide forcing is considered. Velocity (direction and intensity) during (a) high water, (b) the ebb (high water + 3h), (c) low water and (d) during the flood (low water + 3h)

sites, which is also highlighted by the good correlation coefficients between simulated and measured data. At N'Digoro and Platier sites, the simulated currents are clearly out of phase and even nearly in opposition of phase for the Platier site. This phase lags are largely attributed to the role played by the wave-driven component of the cross-reef fluxes which are nearly always present during the field measurements and affect the whole lagoon dynamics (see below) but are not taken into account in the *Reference* simulation. Part of the discrepancy between simulated and measured tidal currents in the N'Digoro passage can also be explained by the fact that this mooring is located in an area with a high spatial variability of current and the spatial resolution of the topography is not enough accurate to precisely reproduce the measured velocity in such a sheared flow at the exact location. Note also that, for the calm wave case, a quarter-diurnal wave at the Ténia site was observed both in numerical and measured data.

The model provides a good representation of the expected periodic filling-emptying cycles of the lagoon induced by the tidal oscillations, see correlations in tables 3 and 4. In most of the lagoon, the currents

reach their maximums during flow and ebb and nearly vanish during slack water, see Fig. 7. Note that, due to the reduced passage section (depth < 3m), strong currents are observed in the strait of La Foa bay. In the studied area (see description in section 2.1), the most intense flows are observed in the main channels of the lagoons. The simulated flow pattern does not show a circulation cell but rather nearly symmetric filling/emptying processes between the south (Ténia passage) and north (Isié openings) connections. The impact of such flow structure in the lagoon and the presence of a nearly "dead" zone in the lagoon center on the water renewal dynamics will be further analysed in section 4.3.

4.1.2 Implementation of cross-reef fluxes

The parameterization of cross-reef fluxes presented in section 3.4 is used to take into account the combined wave and tide driven fluxes above the reef in the numerical model.

In the lagoon numerical model, the depth D over the reef flat is computed at each time step (taking account the water-level and the bathymetry) to determine if incoming waves are expected to break or not on the reef part of the mesh grid. If it does, i.e. for $H_s/H > 0.7$, the cross-reef flux at reef mesh where waves are expected to break is approximated according to the above-mentioned linear adjustment. This value is then imposed to the flux at the corresponding reef mesh cells in the model. When $H_s/H < 0.7$, the fluxes are assumed to be purely driven by tidal oscillations. In this case, no forcing is imposed and the model computation is freely running at this cell.

Note that the water level at the outside reef border grid points is adjusted to compensate the wave-induced flux imposed on the reef. Furthermore, a ramp is applied to ensure a smooth transition between breaking and non breaking states. The imposed flux F is thus defined as:

$$F = F_w + \frac{1 - \tanh(r)}{2} F_c \quad (3)$$

where r is the transition parameter, F_c is the freely calculated flux and F_w the parameterized flux including wave effects given by

$$F_w = H_s(A\eta + B) \frac{1 + \tanh(r)}{2} \quad (4)$$

The transition parameter is defined as $r = \alpha(\frac{H_s}{D} - B_c)$ where α is the ramp steepness (chosen here to be 0.5) and $B_c = 0.7$ is the breaking threshold, see section 3.4 for details.

4.1.3 Analysis of the implementation of cross-reef fluxes

The purpose of this section is to demonstrate how the implementation of cross-reef fluxes improves the numerical description of tide propagation and currents. A more comprehensive analysis of the effect of wave-induced fluxes on the overall lagoon dynamics is proposed in section 4.2.

The comparison between *Reference* simulation and field measurements carried out in section 4.1.1 is based on the M2 tidal component. It is recalled that the M2 component of in-situ tidal currents is extracted from the whole three-months measurement campaign during which the reef is nearly systematically exposed to swell forcing (see Fig. 2). Wave-breaking induced cross-reef fluxes are expected to affect the lagoon dynamics most of the time (see Fig. 6) and thus to largely explain the differences observed between simulated and measured tidal currents observed in table 4, given that the *Reference* simulation does not take into account any wave influence. To test the effect of cross-reef fluxes parameterization, a *Wave* simulation is performed using the averaged value of significant wave height over the whole field data $H_s = 0.66\text{m}$. The effect of parameterization is to shift in time the apex of cross-reef fluxes and, in consequence, also shift the timings of velocity maximums in channels and passages. Table 5 shows the influence of this average wave forcing on the model results, in terms of tidal currents magnitude and phase. The comparison between the *Reference* simulation (4) and the *Wave* simulation (tab. 4) highlights the general improvement of the agreement between model and measurements with increased correlations and lower errors. One notes that the phase lag, previously observed between in-situ measurements and *Reference* simulation (tab. 4), have been significantly reduced.

Another test of the validity of the proposed parameterization is performed on two wave-cases numerical simulations compared with the selected field measurements presented in Fig. 5. The first simulation is a low wave forcing case $H_s=0.3\text{m}$ presented in Fig. 11 (bottom left plot) and to be compared with the low wave field data shown in Fig. 5, left plots. The second simulation represents a significant wave event with $H_s=1.2\text{m}$ (Fig. 11, bottom right plot) related to the high wave measurements presented in Fig. 5, right plots. The primary effect of wave-induced fluxes in the numerical lagoon dynamics is, as described in more

	Magn ^{Mod.} (m/s)	Magn ^{Meas.} (m/s)	Phase ^{Mod.} (°)	Phase ^{Meas.} (°)	Corr. (%)	Error
N'Digoro	0.03	0.08	129	198	36	$3.1 \cdot 10^{-3}$
Isié	0.27	0.29	184	166	95	$4.2 \cdot 10^{-3}$
Platier	0.03	0.03	225	269	72	$0.2 \cdot 10^{-3}$
Ténia	0.15	0.09	178	203	91	$3.0 \cdot 10^{-3}$

Table 5 Comparison between model computations (*Wave* case simulation $H_s = 0.66\text{m}$) and measurements for the M2 tidal component of velocity at each mooring. Current amplitude, phase, correlation and mean squared error.

1.2m wave forcing				
Case	N'Digoro	Isié	Ténia	Platier
<i>Ref.</i> simulation	1	82	-55	34
<i>Wave</i> simulation	92	97	77	67

0.3m wave forcing				
Case	N'Digoro	Isié	Ténia	Platier
<i>Ref.</i> simulation	71	96	35	90
<i>Wave</i> simulation	83	97	51	91

Table 6 Correlation coefficients (in %) between measured and simulated fluxes over two tidal cycles. The measurements, previously presented in Fig. 5, are compared with *Reference* and *Wave* simulations, the latter being forced with the in-situ measured significant wave height.

details in section 4.2, to force water input above the reef barrier and, as a consequence, to enhance outflows in each passages and reef openings (see e.g. Fig. 11, top right plot). One focuses here on the improvements brought by the parameterization in the numerical description of the effect of wave-induced fluxes. It is first recalled that, as only the main tide component is considered (M2 tidal wave), any direct comparison between model and measurement should be interpreted with caution. With this incomplete forcing, the numerical tidal amplitude is lower than the measured one for the selected days and so are the tidal oscillations of cross-reef flux and exchanges through passages and reef-openings. However, accounting for such sources of uncertainties, computation of correlation coefficients between measured and simulated can provide an interesting estimation of the effect of parameterization. The experimental cases presented in section 3.4 (see Fig. 6) are compared with *Reference* and corresponding *Wave* simulations, i.e. $H_s = 0.3$ and 1.2m for the low and high wave cases, respectively. The correlation coefficients between numerical and measured fluxes computed over two successive tidal cycles are shown in table 6. One notes the general improvement in the whole lagoon brought by the use of wave-induced flux parameterization in the numerical model. This is particularly significant for the high wave case ($H_s = 1.2\text{m}$) where the wave effect is expected to be stronger. The greater improvements are observed for the N'Digoro and Platier sites, for which the *Reference* simulation is clearly no able to simulate the fluxes.

This demonstrates, on one hand, the necessity to account for wave effect when performing hydrodynamic simulations of the reef-lagoon system and, on the other hand, the efficiency of the parameterization to represent the main effect of wave-induced flux on the lagoon circulation. Thus, pointing out the process-study objective of the numerical simulations, the present results indicate that the proposed model with its parameterization of wave-induced fluxes is a relevant tool to analyse the response of the reef-lagoon system to wave forcing.

4.2 Cross-reef fluxes impact on lagoon dynamics

4.2.1 Influence of cross-reef tidal fluxes

The first step of the analysis of the role played by cross-reef water fluxes on the lagoon dynamics is to compare the *Reference* and *Wall* simulations. No wave effect is considered and the only difference between test cases is the presence of an impervious reef barrier for the *Wall* simulation, so the sole effect of tidal cross-reef fluxes can be analysed. The comparison of mid-tide currents shown in figures 7 (right plots) and 9 shows that, even if the general circulation pattern within the lagoon and around passages and openings is quite similar for both simulations, an overall flow enhancement is observed when considering the reef as an impervious barrier. This trend for stronger currents is also demonstrated in Fig. 8 which compares the simulated fluxes at each sites over two tidal cycles. Along the tidal cycle, one notes that the difference between *Reference* and *Wall* simulations increases with water level. This trend is expected as the passage

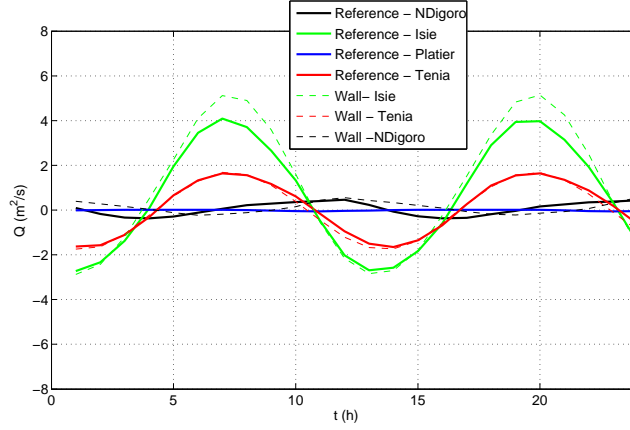


Fig. 8 Influence of cross-reef tidal fluxes: comparison between *Reference* and *Wall* simulations for simulated fluxes over two successive tidal cycles at each site.

	Case	N'Digoro	Isié	Ténia	Platier
Flood	<i>Ref.</i>	-2.1	49.9	49.2	9.3
	<i>Wall</i>	-5.1	64.4	51.6	-
Ebb	<i>Ref.</i>	5.1	-49.7	-42.7	-20.8
	<i>Wall</i>	12.1	-67.2	-52.5	-
Total	<i>Ref.</i>	3	0.3	6.6	-11.6
	<i>Wall</i>	7.1	-2.6	-0.77	-

Table 7 Exchanged volumes (in millions of m^3) through N'Digoro and Ténia passages and Isié reef opening for the *Reference* and *Wall* simulations. Calculations are performed for flood and ebb tides and for complete tidal cycle.

section above the reef is maximal at high tide. Most of changes in terms of magnitude are observed in the Isié site, which is explained by the fact that, in the *Wall* case, the Isié reef opening becomes the only direct connection with open ocean in the studied area. A shift in phase is observed in the N'Digoro site: the internal exchanges processes within the lagoon due to filling at flood tide are delayed in the *Wall* simulation as the input section for ocean waters is reduced.

The exchanged water volumes are also strongly enhanced by the presence of a wall reef (see table 7), in particular in the Isié site (up to 35% during ebb tide) and through the N'Digoro passage (up to 137% during flood tide). This results in a sign inversion of the total volume exchanged at Isié and Ténia sites: the volume balance being a net input for the *Reference* simulation becomes a net output in the *Wall* case. In addition, computations of exchanged volume above the barrier for the complete total cycle are 1.8 to 3.9 times the volumes exchanged at the other sites (see table 7). It is thus demonstrated that, the tidal cross-reef flows play a significant role in the lagoon-reef system dynamics and must be taken into account for any hydrodynamic study of the channel lagoons as the one considered here.

4.2.2 Influence of wave breaking induced cross-reef fluxes

This section is dedicated to the numerical study of the impact of wave-induced cross-reef water fluxes on lagoon dynamics. The two *Wave* simulations ($H_s = 0.3$ and 1.2 m) presented in section 4.1.3 are used to characterize the effect of swell forcing on the lagoon dynamics. A first qualitative analysis is performed in Fig. 10 which compares the lagoon circulation patterns simulated during flood and ebb tides for both *Wave* simulations. In the low wave ($H_s = 0.3$ m) case (Fig. 10, top plots), the effect of wave breaking, which only occurs for nearly mid to low tide, is restricted to areas close to the reef barrier. One notes that, at the selected times, the flow over the reef is almost systematically inward and does not reverse during the ebb tide as observed for the *Reference* simulation (Fig. 7, top right plot). More complex vortex-like flow patterns can be observed in the sheared zones surrounding the reef openings (Fig. 10, top left plot). In the high wave case (Fig. 10, bottom plots), the effect of the parameterized wave-induced fluxes above the barrier is much more straightforward. During the ebb tide, the additional fluxes caused by wave breaking all along the reef flat force water inputs above the barrier and, as a consequence, push out water in each of the reef openings and passages. During the flood, the cross-reef fluxes above the barrier participate in the

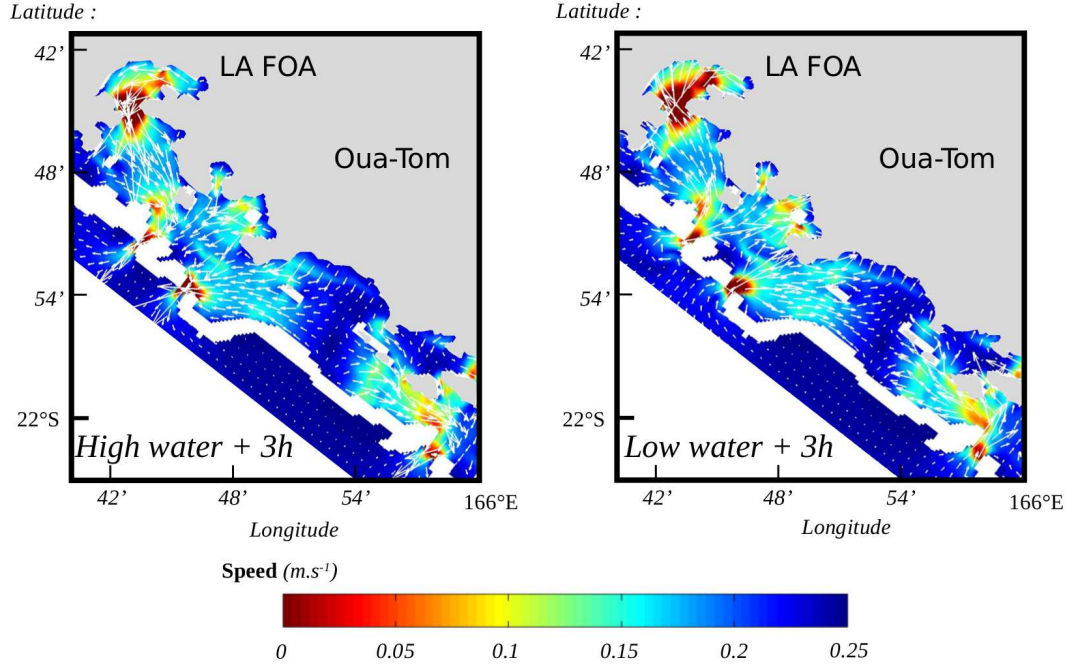


Fig. 9 Tidal current (during ebb and flood) in case of a wall instead of a reef barrier. Velocity (direction and intensity) during (a) the ebb (high water + 3h) and (b) during the flood (low water + 3h)

lagoon filling, reduce the surface gradients between interior lagoon and open ocean and thus weaken the flow through Isié and Ténia passages.

Quantitative impact of wave-driven flux parameterization on the lagoon dynamics is shown in Fig. 11 which compares the simulated fluxes through reef opening and passages and over the reef for the *Reference* simulation and both *Wave* cases. The effect is rather small for the low wave case ($H_s = 0.3\text{m}$) and mostly observed around low tide for small H_s/H ratio. Maximal outflows observed just after low tide at Isié, Ténia and N'Digoro sites are increased by 12, 21 and 57 % in comparison with the *Reference* simulation, respectively. The absolute value of cross-reef flows remains very weak both in *Reference* and 0.3m *Wave* simulations but its integrated effects along the whole barrier is significant on the exchanged volumes (see below). The modifications induced in the 1.2m *Wave* simulation are much more significant. For such incoming wave height, the breaking threshold is exceeded all along the tidal cycle. This induces, (i), a systematically inward positive flow above the reef, (ii), a permanently outflow leaving the lagoon in the N'Digoro passage and, (iii) a shift toward negative values of tide-related oscillations of fluxes calculated in Isié and Ténia sites. The increases of maximal outflow for Isié, Ténia and N'Digoro sites in comparison to the *Reference* simulation are 71, 104 and 346 % respectively.

An ensemble of *Wave* simulations has been performed to quantify the dependency of exchanged volumes between lagoon and open ocean at each site on the incoming wave height. The results, shown for ebb, flood and total tidal cycles in Fig 12, confirm the observations made hereinbefore. One notes first that the parameterization of wave-induced fluxes starts to affect the exchanged volumes from incoming wave height around 0.2m . The main trend when wave forcing increases is an increasing input (positive volume) of water over the reef coupled with an increasing output (negative volume) at other sites. For each site, one notes the presence of threshold in term of wave height from which a reversal in flow pattern is observed. For instance, during the flood, the water inputs observed for the *Reference* simulation at Isié and Ténia are reversed toward a net output of water around $H_s = 1.6$ and 1.2m , respectively. During the ebb, the initial input (resp. output) observed at N'Digoro and Platier sites are reversed into output (resp. input) for H_s around 0.3m . The observed trends are confirmed when computing the total volumes exchanged over a tidal cycle: the cross-reef fluxes induced by wave breaking over the reef enhance the emptying processes at Ténia, Isié and N'Digoro sites.

These exchanged volumes can be compared to the total average volume of the considered lagoon area ($1.3 \cdot 10^9 \text{m}^3$). Table 8 shows the relative exchanged volume over a complete tidal cycle for the *Reference* and two *Wave* simulations with $H_s = 0.6$ and 1.2m . This demonstrates the importance of wave-induced

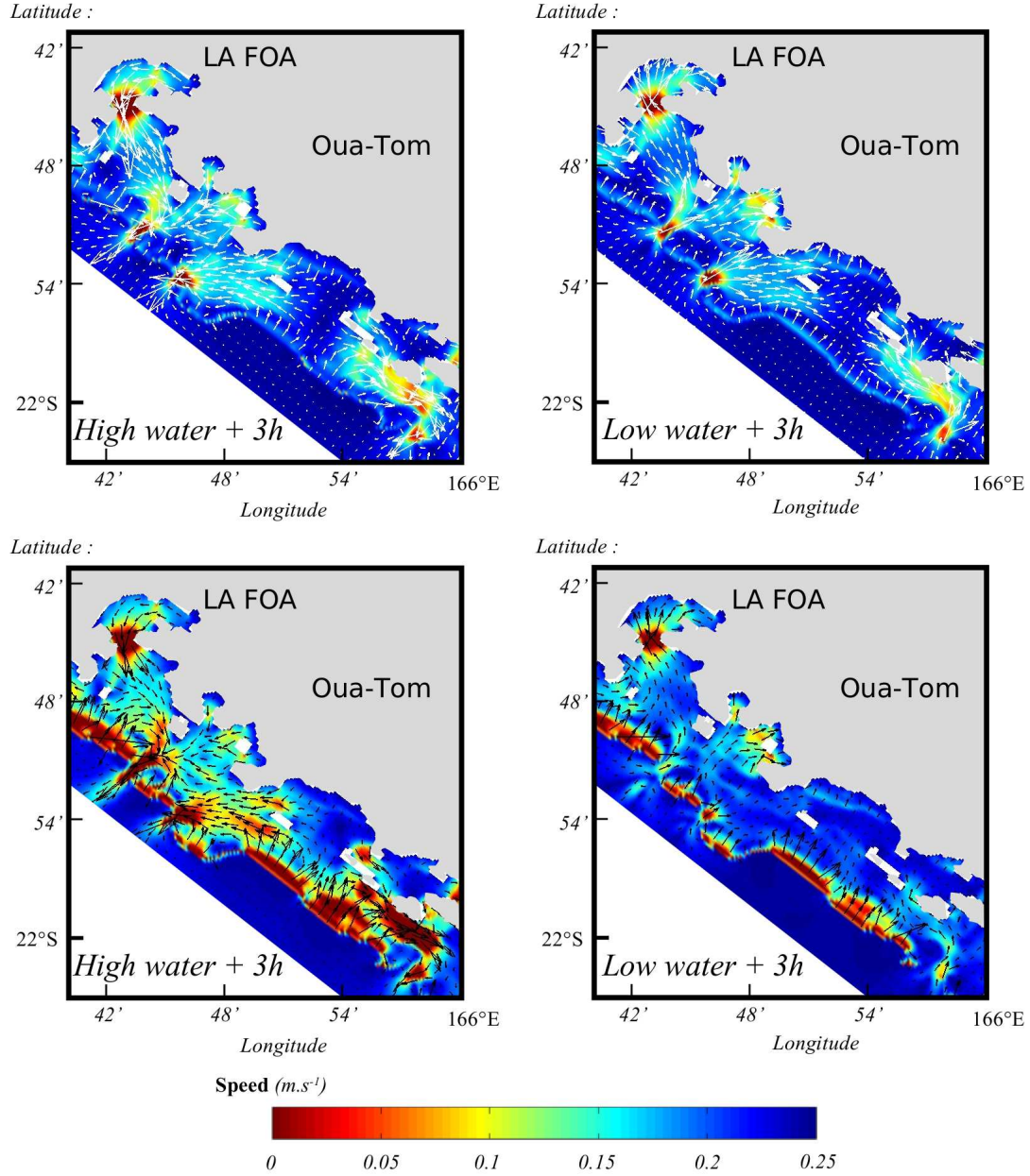


Fig. 10 Numerical lagoon dynamics for the 0.3m wave case (top plots) and 1.2m wave case (bottom plots). Velocity (direction and intensity) during ebb (high water + 3h, left plots) and during the flood (low water + 3h, right plots)

cross-reef fluxes in the exchanges between inner lagoon and open ocean. By the sole effect of tide (*Reference* simulation), the exchanged volumes integrated over a complete tide cycle at each site represent less than 1% of the lagoon volume. As wave height increases, this ratio becomes more and more significant, up to 9.4 and 18.75% for the 1.2m *Wave* simulation at Isié and Platier sites, respectively. For the *Reference* simulation, the tide-averaged volumes are moving into the lagoon at Isié, N'Digoro and Ténia sites and toward the ocean above the barrier at the Platier site. For the *Wave* simulations, the directions are totally reversed: inflow over the reef barrier and outflows at other sites. The direction as well as the order of magnitude of exchanged volumes are in much better agreement with estimations from in-situ measurements (see section 3.3.2) when considering wave effect in the model. The major contribution of wave-induced fluxes implies important consequences on the age of lagoon waters which are analysed in the next section.

In addition, it has been observed that the tidal dynamics inside the lagoon is also affected by the wave-driven fluxes above the barrier. Indeed, although amplitude of the tidal current in the lagoon are weakly

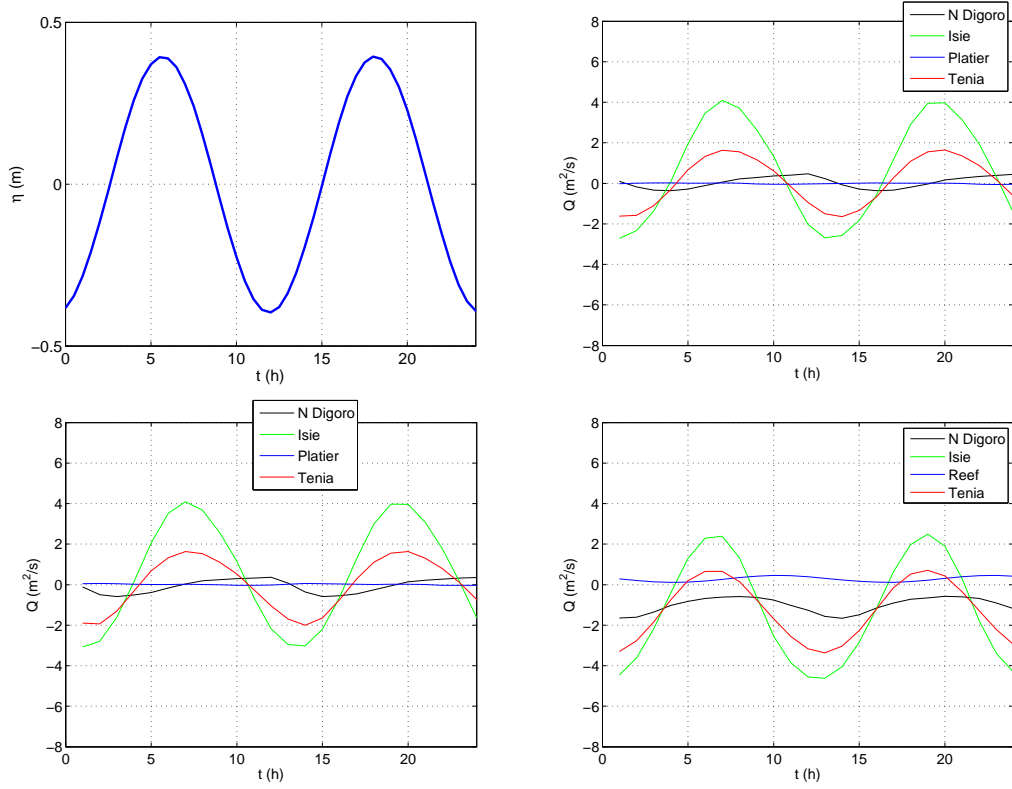


Fig. 11 Impact of wave-driven flux parameterization on the lagoon dynamics simulation for two tidal cycles. Left top: mean water level. Right top, left bottom and right bottom: Numerical flux through reef opening and passages and over the reef for the no-wave case, 0.3m wave case and 1.2m wave case, respectively. Positive velocity values correspond to inward current entering the lagoon.

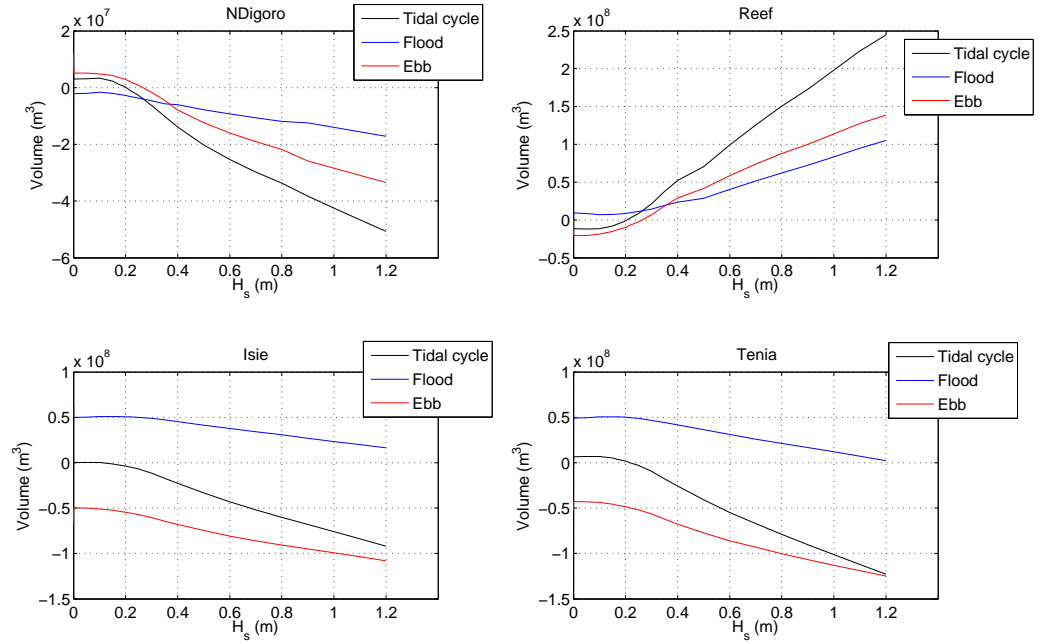


Fig. 12 Exchanged volumes during flood, ebb and complete tidal cycles versus incoming wave height. Note the scale difference for the y-axis.

	N'Digoro	Isié	Ténia	Platier
<i>Ref.</i>	0.23	0.02	0.46	-0.85
<i>Wave</i> $H_s = 0.6\text{m}$	-1.9	-3.3	-4.2	7.6
<i>Wave</i> $H_s = 1.2\text{m}$	-3.9	-7.0	-9.4	18.5

Table 8 Tidally averaged exchanged volumes (in % of the total lagoon volume) through N'Digoro and Ténia passages, Isié reef opening and over the reef Barrier for the *Reference* and two *Wave* simulations with $H_s = 0.6$ and 1.2m . Relative volumes are computed over a complete tidal cycle.

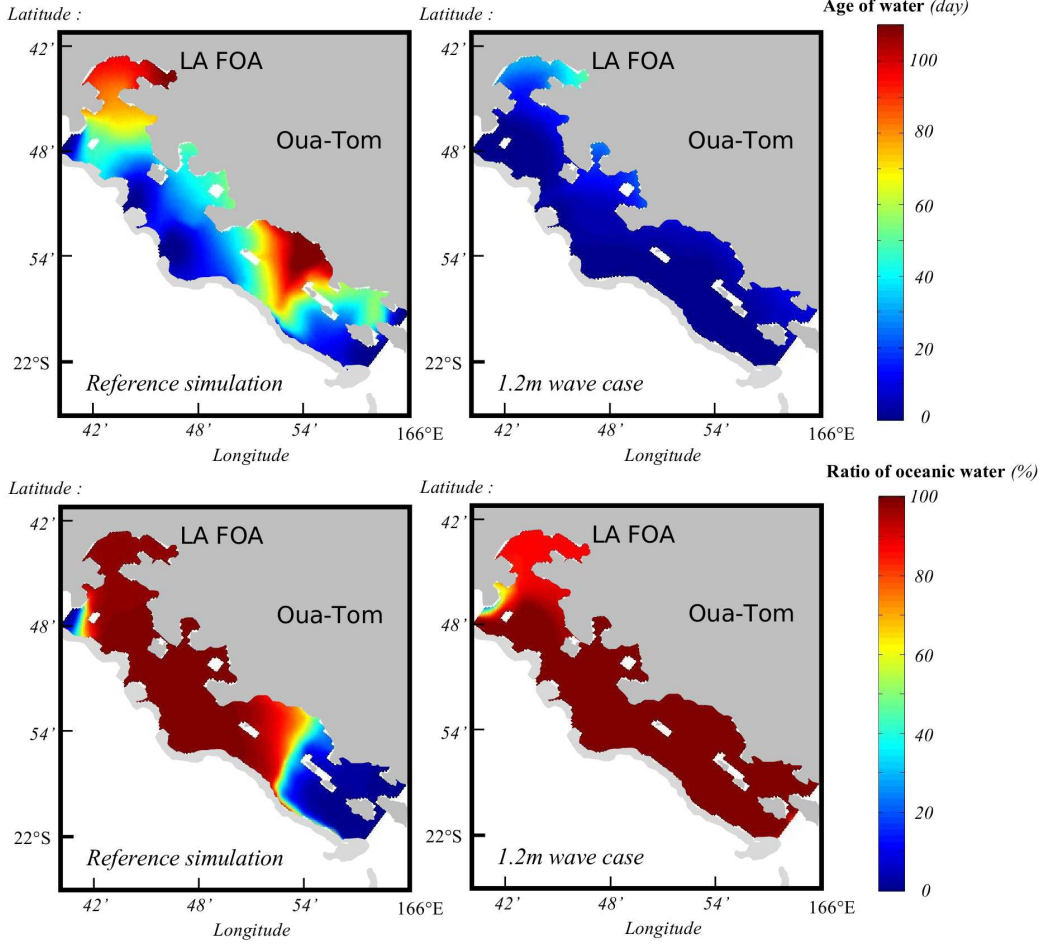


Fig. 13 Spatial distributions of lagoon water age (top plots) and fraction of oceanic water (bottom plots) for the *Reference* and 1.2m *Wave* simulations in left and right plots, respectively.

impacted, the phases of the current are strongly modified. For instance, in most of the lagoon area, a time lag greater than 1 hour can be observed between the current maximums simulated in the no-wave and 1.2m wave cases. This time lag appeared even in bays along the coast and is accentuated close to the reef.

4.3 Age and origin of lagoon water

The action of cross-reef fluxes on the overall circulation naturally affects the water distribution and retention in the lagoon. Based on the methods referred to in section 2.3.6, computations of age and origin of the lagoon water are shown in Fig. 13 for the *Reference* and 1.2m *Wave* simulations. Both spatial distribution and magnitude of the water age in the lagoon (Fig. 13, top plots) are strongly affected by the presence of wave. In the *Reference* simulation, the most recent waters are observed in the vicinity of passages and openings. Further inside the lagoon, the age becomes rapidly greater than 40 days and reaches more than 100 days in two specific areas: the La Foa bay and the central southern lagoon. The La Foa bay is the

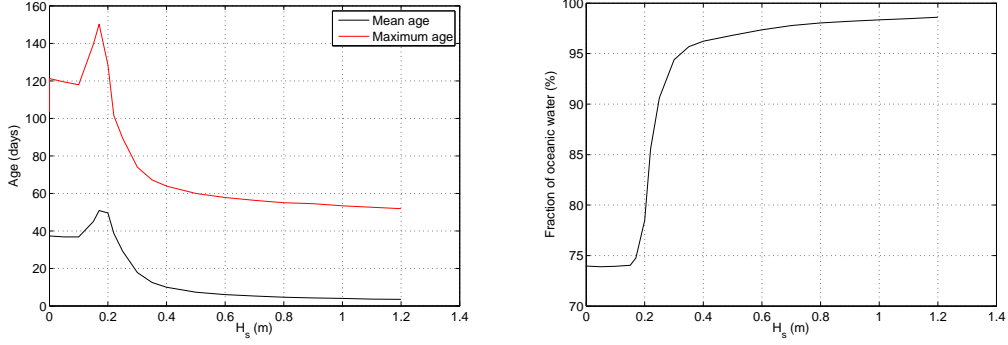


Fig. 14 Age (averaged and maximal values over the whole lagoon) and fraction of oceanic water versus incoming wave height in left and right plots, respectively.

most remote place of the studied lagoon, separated by a relatively narrow and shallow strait, and, as such, is more protected for water renewal processes. The important water age values observed in the central southern lagoon are related to the hydrodynamic features described in section 4.1.1 (see Fig. 7). When only forced by the tidal oscillations of the free surface, a nearly symmetric pattern of lagoon filling/emptying cycles is observed between Ténia and Isié. This leads to the presence of a nearly "dead" zone within which the water age varies from about 60 days near the reef barrier to 110 days close to the mainland shore. The *Wave* case shows a radically different configuration. The maximal age is also observed at the end of the La Foa bay, but do not exceed 60 days and is even less than 30 days in most of the bay. In the main lagoon, the age is generally less than one week. The "dead" zone is no more observed and the age displays a nearly cross-shore distribution being less than one day near the reef and increasing up to 10 days at the shoreline and about 20 days in the protected areas.

The bottom plots of Fig. 13 present the spatial distribution of the fraction of oceanic waters in the lagoon for the *Reference* and 1.2m *Wave* simulations. One recalls here that the computation is initiated assuming three distinct water masses: the lagoon itself, the neighboring lagoons and the open ocean. The procedure aims to discriminate the spatial origins of the studied lagoon waters depending on the incoming wave conditions. The difference between *Reference* and 1.2m *Wave* simulations is straightforward. In the absence of waves, the central and northern areas of the lagoon are mostly filled with oceanic waters while southern lagoon waters have been supplied by the neighboring lagoon through the Ténia passage. When waves break over the reef, the cross-reef input of open ocean water is observed to fill nearly the entire lagoon with outside oceanic waters except in the La Foa bay where the proportion of water coming from northern lagoon increases with wave height.

Figure 14 depicts the evolution of lagoon water age (left plot) and fraction of oceanic water (right plot) under variable wave forcing. The age plot shows, as expected, an overall decrease for both mean and maximal values computed over the studied area when wave height increases. The mean value of lagoon water age, is estimated to be around 35 days for the *Reference* simulation and reaches a nearly asymptotic value of about 3 days for large waves ($H_s > 0.4$ m). The maximal age starts from 120 days in the no-wave case to 50 days for $H_s = 1.2$ m. An increase of the age of lagoon water is observed for very small wave ($0.1 < H_s < 0.2$ m). This peculiar feature is a transitional behavior explained by the fact that, for such small wave forcing, the wave-induced fluxes above the reef are sufficient to push water toward the coast and to increase the retention of inner lagoon waters in particular in the northern/central part but is not strong enough to tip the hydrodynamic system in the typical *Wave* configuration observed previously. Figure 14 (right plot) shows that, in the no-wave case, nearly 74% of lagoon water comes from the open ocean and the rest originates from neighboring lagoons. As soon as the wave height reaches 0.4m, more than 95 % of the studied water volume comes from the ocean. This proportion then shows an asymptotic behavior when the wave height increases (98.6 % for $H_s = 1.2$ m).

Computations of age and origin of the lagoon water demonstrate the flushing effect of wave forcing on the lagoon circulation, which is expected to play a preponderant role on water quality and bio-chemical dynamics of the whole reef-lagoon system. A very interesting observation in figure 14 is that most of the modifications of lagoon water properties are observed for H_s ranging between 0.2 and 0.4m. Under 0.2m wave height, the lagoon functioning is mainly controlled by the tidal filling and emptying while above $H_s = 0.4$ m, the lagoon functioning tends toward an asymptote.

5 Conclusion and prospects

This paper reports on a combined experimental/numerical study of the impact of cross-reef fluxes on the hydrodynamics of a channel reef-lagoon system. The selected study site is the Ouano lagoon, New Caledonia, France. A three months survey of wave transformation over the reef barrier and water fluxes through main passages and reef opening has been carried out during the austral winter season.

An overall processing of the available field data revealed the driving factors of the lagoon dynamics. It is observed that the circulation at the lagoon scale and the water fluxes through the passages are primarily driven, as expected, by the tidal oscillations. The offshore wave forcing, in particular wave height, is observed to play a secondary but significant role. Influence of other factors, such as incident swell wavelength and angle, wind speed and direction or rainfall, has been observed to be much less significant for the studied period.

Depth-averaged currents above the reef are generally smaller than their counterparts through passages and reef openings but, due to the very wide area of the reef flat, they contribute significantly to the total exchanges between lagoon and open sea. These cross-reef fluxes are induced by the combined action of tides and breaking of incoming offshore waves over the reef slope and flat. An ad-hoc breaking threshold is introduced, relating the significant wave height to the water depth over the reef top. It is observed that, as soon as the breaking threshold is reached, the measured cross-reef velocity increases with increasing swell wave height and mean water level. The enhanced water input above the reef observed during large swell events is balanced by outflows in each passages and reef opening. The estimate of wave-induced cross-reef fluxes is thus of primary importance when simulating the overall lagoon circulation.

Based on the field data recovered during the three months survey of the lagoon, a simple parametrization of the cross-reef fluxes is proposed and implemented in a coastal lagoon circulation model (ROMS). The parametrization clearly improves the quality of the model description in wave-breaking conditions and emphasizes the importance of cross-reef currents in such a narrow lagoon dynamics. The fluxes exchanged in reef openings and in passage depend clearly on the wave height. The lagoon functioning (age and origin of lagoon water) is also observed to depend on the wave height with the presence of thresholds occurring for 0.2 and 0.4m wave height. Under and above these limits, the lagoon functioning varies weakly with wave height, while between these two values, strong modifications are observed.

Further research prospects include the progressive implementation of additional physical processes (wind, rainfall, stratification, etc) into the model, the extension of the in-situ survey of lagoon dynamics to other seasons and wave and wind climate conditions, the refining of the parameterization, in particular to account for incoming wave direction, and, to finally allow the model validation in an operational point of view, a set of currents and water properties measurements within the lagoon. In a longer term inter-disciplinary approach, one guiding prospect of our work is the comparison with biological field data of plankton dynamics. In particular, this hydrodynamics model may be coupled with a biological model in order to understand the importance of cross-reef fluxes on the biology of the inner lagoon and the filtering role of the coral reef which consumes the picoplankton, and may induce an increase of the amount of food directly consumable by benthic reef community [52,53].

Site		ADCP							Pressure sensors	
Location	Depth	Sensor	Samp.-Aver.	Vert. res.	1st Bin	Pings/ens	Time/Ping	SD	Sensor	Integration time
	(m)	(min - min)	(m)	(m)			(s)	(cm/s)		(min)
N Digoro	14.84 \pm 0.05	Rdi 300KHz	10-1	1.5	3.73	100	0.55	1	Micrel SP2T100	10
Isie	16.80 \pm 0.05	Rdi 300KHz	10-1	1.5	3.73	100	0.55	1	ADCP	10
Platier	3.53 \pm 0.05	Sontek 3MHz	10-1.5	0.25	0.45	asap	nc	nc	Micrel SP2T100	10
Ténia	13.89 \pm 0.05	Rdi 300KHz	10-1	1.5	3.73	100	0.55	1	ADCP	10

Table 9 Sensor characteristics during the field experiment (*asap*: as soon as possible - ping per ensemble, *nc*: no theoretical value calculated by Sontek).

			M2		S2		K1		M4		Ratio
			Value	Error	Value	Error	Value	Error	Value	Error	
N Digoro	Water level	Amplitude (m)	0.38	0.004	0.16	0.003	0.13	0.004	0	-	91%
		Phase ($^{\circ}$)	272	0.5	288	1.5	14	1.8	-	-	
		SNR	9.10 ³		2.10 ³		1.10 ³		0		
	Velocity	Amplitude (m/s)	0.08	0.002	0.04	0.002	0.06	0.002	0.01	0.002	66%
		Phase ($^{\circ}$)	198	1.4	227	3.2	268	2.4	231	10.7	-
		SNR	2.10 ³		6.10 ²		5.10 ²		31		-
Isie	Water level	Amplitude (m)	0.36	0.002	0.15	0.002	0.13	0.002	0	-	97%
		Phase ($^{\circ}$)	272	0.2	289	0.6	17	0.7	-	-	
		SNR	5.10 ⁴		8.10 ³		6.10 ³		5		
	Velocity	Amplitude (m/s)	0.29	0.002	0.11	0.002	0.07	0.002	0.03	0.02	93%
		Phase ($^{\circ}$)	166	0.4	191	1.1	255	1.9	356	3.9	
		SNR	2.10 ⁴		3.10 ³		1.10 ³		3.10 ²		
Platier	Water level	Amplitude (m)	0.4	0.002	0.18	0.002	0.11	0.001	0	-	97%
		Phase ($^{\circ}$)	272	0.2	293	0.5	23	0.8	-	-	
		SNR	7.10 ⁴		2.10 ⁴		5.10 ³		3		
	Velocity	Amplitude (m/s)	0.03	0.001	0.01	0.001	0.02	0.001	0.01	0.001	58%
		Phase ($^{\circ}$)	269	1.2	293	5	34	3	24	6	
		SNR	1.10 ³		4.10 ²		1.10 ²		1.10 ²		
Ténia	Water level	Amplitude (m)	0.41	0.001	0.17	0.001	0.13	0.002	0	-	97%
		Phase ($^{\circ}$)	272	0.2	292	0.5	17	0.7	-	-	
		SNR	1.10 ⁵		1.10 ⁴		1.10 ⁴		2		
	Velocity	Amplitude (m/s)	0.09	0.002	0.04	0.002	0.02	0.002	0.01	0.001	67%
		Phase ($^{\circ}$)	203	1.3	195	2.6	20	6	229	10	
		SNR	3.10 ³		6.10 ²		99		35		

Table 10 Tidal components of water level and velocity along the main axis established using Foreman spectral tidal analysis [49]. Ratio: Variability explained by the harmonic analysis, ratio between inferred function and data (in percent). SNR: signal-to-noise power ratio. It is computed based on the square of the ratio of amplitude to amplitude error. This parameter indicates the degree of confident of this discretization [50]: high level of confident for SNR>10, and not bad for SNR as low as 2 or 3. Grey values represent values which are not accurate due a too short duration of the measurements with respect to the Rayleigh criteria (mooring S1) or due to the low SNR.

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